

Evaluation of vertical mass flux in high-resolution

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Motivation

Craig and Cohen (2006) used the Gibbs canonical ensemble from statistical mechanics to derive equilibrium statistics of a field of cumulus clouds under homogeneous large-scale forcing.

They derived the probability density function of individual cloud mass fluxes, in the limit of non-interacting convective cells, to be exponential:

$$p(m)dm = \frac{1}{\langle m \rangle} e^{-m/\langle m \rangle} dm$$

Mean cloud number density distribution:

Validation of their theory with CRM simulations at 2 km horizontal resolution in radiative convective equilibrium for different forcings :

Model description

Simulations are performed with the anelastic, nonhydrostatic model EULAG (<u>EU</u>lerian/semi-<u>LAG</u>rangian fluid solver), which uses a second order accurate, semi-implicit, non-oscillatory forward in time approach (Prusa et al., 2008).

The underlying transport operator to solve the moist governing equations is formulated for arbitrary curve-linear frameworks and employs the MPDATA (multidimensional positive, definite advection transport) algorithm.







 $\langle N \rangle$: Ensemble mean number of clouds $\langle m \rangle$: Ensemble mean mass flux per cloud

- → exponential distribution is independent of the forcing
- → increasing the forcing mainly increases the number of clouds, not cloud strength

Is this theory of an exponential mean cloud number density distribution still valid at very high horizontal resolutions (~100 m), where small cumulus clouds are actually resolved? The buoyancy term is expressed by the perturbation of the density potential temperature:

 $\theta_d' = \theta + \overline{\theta} \left(\epsilon \ q_v - q_c - q_p \right)$

(Mixing ratio of water vapor (q_v), cloud water (q_c) and rain water (q_p))



Control simulation: set-up

- 3D model domain: 128 km *128 km * 20 km
- 2 km horizontal, 200 m vertical resolution
- Sea Surface Temperature fixed at $\Theta_{surf} = 300 \text{ K}$
- Periodic boundary conditions in x and y
- Rayleigh damping layer at top of the domain
- Horizontally homogeneous radiative cooling rate F_{rad}





Evolution to quasi-equilibrium

The model is run with $F_{rad} = -8$ K/day from an initial horizontally homogeneous state with no convection:

- → high frequency variability (~1 h) can be directly related to convective activity
- → long, slow trend of the overall spin-up towards radiative convective equilibrium (~30 days)

In equilibrium, the input of energy into the system (surface fluxes) provides exactly the energy required by convection to offset F_{rad} .



Exponential distribution of mass flux per cloud





The equilibrium state



- 1) Criterion to define cloudy grid points: w > 1 m/s and q_c > 1.e-3 g/kg
- 2) Search for adjacent cloudy grid points
- 3) Compute mass flux per cloud: $m_i = \rho * \sigma_i * < w_i > c$

(σ_i :size of the cloud, ρ :density of air, $\langle w_i \rangle$:average vertical velocity)



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